The pdf of Sea Surface Winds: Effects of Variable Stratification and Boundary Layer Height

Adam Monahan

monahana@uvic.ca

School of Earth and Ocean Sciences, University of Victoria Victoria, BC, Canada





The pdf of Sea Surface Winds: Effects of Variable Stratification and Boundary Layer Height – p. 1/18

Surface Exchange



Surface Exchange

surface fluxes depend on surface winds, in general nonlinearly



- Surface Exchange
 - surface fluxes depend on surface winds, in general nonlinearly
 - \Rightarrow mean fluxes depend on higher-order wind moments



Surface Exchange

surface fluxes depend on surface winds, in general nonlinearly

 \Rightarrow mean fluxes depend on higher-order wind moments

Power Generation



Surface Exchange

- surface fluxes depend on surface winds, in general nonlinearly
- \Rightarrow mean fluxes depend on higher-order wind moments
- Power Generation
 - wind power potentially significant source of energy



Surface Exchange

- surface fluxes depend on surface winds, in general nonlinearly
- \Rightarrow mean fluxes depend on higher-order wind moments
- Power Generation
 - wind power potentially significant source of energy
 - wind power density scales as cube of wind speed; higher-order moments important



Surface Exchange

- surface fluxes depend on surface winds, in general nonlinearly
- \Rightarrow mean fluxes depend on higher-order wind moments
- Power Generation
 - wind power potentially significant source of energy
 - wind power density scales as cube of wind speed; higher-order moments important

Extremes



Surface Exchange

- surface fluxes depend on surface winds, in general nonlinearly
- \Rightarrow mean fluxes depend on higher-order wind moments
- Power Generation
 - wind power potentially significant source of energy
 - wind power density scales as cube of wind speed; higher-order moments important

Extremes

extreme high winds have important social, economic impacts



Surface Exchange

- surface fluxes depend on surface winds, in general nonlinearly
- \Rightarrow mean fluxes depend on higher-order wind moments
- Power Generation
 - wind power potentially significant source of energy
 - wind power density scales as cube of wind speed; higher-order moments important
- *Extremes*
 - extreme high winds have important social, economic impacts
- This talk will consider influence of variable surface stratification and boundary layer thickness on the wind speed pdf



Vector Wind Moments



The pdf of Sea Surface Winds: Effects of Variable Stratification and Boundary Laver Height – p. 3/18

Mean and Skewness of Vector Wind

Joint pdfs of mean and skew for zonal and meridional winds





Wind Speed Moments





The pdf of Sea Surface Winds: Effects of Variable Stratification and Boundary Laver Height – p. 5/18

The pdf of wind speed w has traditionally (and empirically) been represented by 2-parameter Weibull distribution:



$$p(w) = \frac{b}{a} \left(\frac{w}{a}\right)^{b-1} \exp\left[-\left(\frac{w}{a}\right)^{b}\right]$$

skew(w) = S(b)
b \approx mean(w)/std(w)



The pdf of wind speed w has traditionally (and empirically) been represented by 2-parameter Weibull distribution:



a is the scale parameter (pdf centre)



The pdf of wind speed w has traditionally (and empirically) been represented by 2-parameter Weibull distribution:



$$p(w) = \frac{b}{a} \left(\frac{w}{a}\right)^{b-1} \exp\left[-\left(\frac{w}{a}\right)^{b}\right]$$

skew(w) = S(b)
b \simeq mean(w)/std(w)

a is the scale parameter (pdf centre)

b is the shape parameter (pdf tilt)



The pdf of wind speed w has traditionally (and empirically) been represented by 2-parameter Weibull distribution:



$$p(w) = \frac{b}{a} \left(\frac{w}{a}\right)^{b-1} \exp\left[-\left(\frac{w}{a}\right)^{b}\right]$$

skew(w) = S(b)
$$b \simeq \operatorname{mean}(w)/\operatorname{std}(w)$$

a is the <u>scale</u> parameter (pdf centre)

- *b* is the shape parameter (pdf tilt)
- $\square p_w(w)$ is unimodal



Wind speed pdfs: observed and modelled



Black line: Weibull, Red dots: idealised BL momentum budget model



- u = along-mean wind component
- v = cross-mean wind component

$$\mathbf{u} = (u, v)$$

$$w = \text{wind speed}$$

h =boundary layer depth

$$T = SAT-SST$$



Mechanistic model of p(w) based on boundary layer momentum budget



- Mechanistic model of p(w) based on boundary layer momentum budget
- Previous work considered combined influence of



- Mechanistic model of p(w) based on boundary layer momentum budget
- Previous work considered combined influence of
 - surface drag (nonlinear)



- Mechanistic model of p(w) based on boundary layer momentum budget
- Previous work considered combined influence of
 - surface drag (nonlinear)
 - downwards mixing of momentum from aloft (constant)



- Mechanistic model of p(w) based on boundary layer momentum budget
- Previous work considered combined influence of
 - surface drag (nonlinear)
 - downwards mixing of momentum from aloft (constant)
 - "ageostrophic forcing" (mean and fluctuating)



- Mechanistic model of p(w) based on boundary layer momentum budget
- Previous work considered combined influence of
 - surface drag (nonlinear)
 - downwards mixing of momentum from aloft (constant)
 - "ageostrophic forcing" (mean and fluctuating)
- Momentum budget averaged over BL (neglecting horiz. advection)

$$\frac{du}{dt} = \underbrace{\overbrace{\overline{U}_s}^{\mathbf{A}}}_{\tau_s} + \underbrace{\overbrace{\eta_u}^{\mathbf{B}}}_{\tau_s} - \underbrace{\overbrace{c_d(w,T)}^{\mathbf{C}}wu}_{h} + \underbrace{\overbrace{w_e}^{\mathbf{D}}(U(h+\delta)-u)}_{h} + \underbrace{\sigma_u\dot{W_1}}_{\sigma_u\dot{W_1}}$$
$$\frac{dv}{dt} = \underbrace{\frac{\eta_v}{\tau_s} - \frac{c_d(w,T)}{h}wv}_{\tau_s} + \frac{w_e}{h}(V(h+\delta)-v) + \sigma_u\dot{W_2}$$
/ic

The pdf of Sea Surface Winds: Effects of Variable Stratification and Boundary Layer Height – p. 9/18

Earlier studies assumed



- Earlier studies assumed
 - neutral surface
 - stratification



- Earlier studies assumed
 - neutral surface
 - stratification
 - \blacksquare constant h



- Earlier studies assumed
 - neutral surface stratification
 - \blacksquare constant h
- Non-neutral surface
 stratification influences drag
 coefficient



- Earlier studies assumed
 - neutral surface stratification
 - \blacksquare constant h
- Non-neutral surface
 stratification influences drag
 coefficient
 - unstable stratification

 $\Rightarrow c_d \uparrow$



- Earlier studies assumed
 - neutral surface stratification
 - \blacksquare constant h
- Non-neutral surface
 stratification influences drag
 coefficient
 - unstable stratification

 $\Rightarrow c_d \uparrow$

stable stratification

 $\Rightarrow c_d \downarrow$



- Earlier studies assumed
 - neutral surface stratification
 - \blacksquare constant h
- Non-neutral surface
 stratification influences drag
 coefficient
 - unstable stratification

$$\Rightarrow c_d \uparrow$$

- stable stratification
 - $\Rightarrow c_d \downarrow$



COARE Parameterisation



Wind variability on shorter timescale than T



- Wind variability on shorter timescale than T
- Sensible to think of w pdf equilibrated to instantaneous T: p(w|T)



- Wind variability on shorter timescale than T
- Sensible to think of w pdf equilibrated to instantaneous T: p(w|T)

Then have

$$p(w) = \int p(w|T)p(T) \ dT$$



- Wind variability on shorter timescale than T
- Sensible to think of w pdf equilibrated to instantaneous T: p(w|T)

Then have

$$p(w) = \int p(w|T)p(T) \ dT$$

• h constant \Rightarrow analytic solution of Fokker-Planck equation for p(w|T)



- Wind variability on shorter timescale than T
- Sensible to think of w pdf equilibrated to instantaneous T: p(w|T)

Then have

$$p(w) = \int p(w|T)p(T) \ dT$$

• h constant \Rightarrow analytic solution of Fokker-Planck equation for p(w|T)

For simplicity, take p(T) Gaussian

 $p(T) \sim \mathcal{N}(\mu_T, \sigma_T)$



- Wind variability on shorter timescale than T
- Sensible to think of w pdf equilibrated to instantaneous T: p(w|T)

Then have

$$p(w) = \int p(w|T)p(T) \ dT$$

• h constant \Rightarrow analytic solution of Fokker-Planck equation for p(w|T)

For simplicity, take p(T) Gaussian

$$p(T) \sim \mathcal{N}(\mu_T, \sigma_T)$$

■ p(w) computed over realistic range of μ_T , σ_T



- Wind variability on shorter timescale than T
- Sensible to think of w pdf equilibrated to instantaneous T: p(w|T)

Then have

$$p(w) = \int p(w|T)p(T) \ dT$$

• h constant \Rightarrow analytic solution of Fokker-Planck equation for p(w|T)

For simplicity, take p(T) Gaussian

$$p(T) \sim \mathcal{N}(\mu_T, \sigma_T)$$

- p(w) computed over realistic range of μ_T , σ_T
- Surface stratification effects \Rightarrow insignificant influence on shape of p(w)





Wind stress - SST coupling: observations



Fig. 2. Maps of spatially high-pass filtered 2 months (May-June 2003) average wind stress magnitude (N m⁻², color) and SST (°C, contours, interval 0.5 °C, zero contour omitted). Data from QuikSCAT scatterometer and AMSR-E. (a) North-west Pacific, Kuroshio region, (b) North-west Atlantic, Gulf Stream and North Atlantic Current region, (c) South-west Atlantic, Brazil-Malvinas confluence, and (d) Southern Indian Ocean, Agulhas Return Current.

Warm SST \Rightarrow *faster* winds (e.g. Small et al., Dyn. Atmos. Oceans, 2008)

The pdf of Sea Surface Winds: Effects of Variable Stratification and Boundary Layer Height – p. 13/18

 Marine boundary layer thickness variable, driven by surface and BL top buoyancy fluxes, shear



Marine boundary layer
 thickness variable, driven by
 surface and BL top buoyancy
 fluxes, shear

Variable boundary layer depth



- Marine boundary layer
 thickness variable, driven by
 surface and BL top buoyancy
 fluxes, shear
- Variable boundary layer depth
- ⇒ "intermittent" downwards
 mixing of momnetum in
 growing BL



- Marine boundary layer
 thickness variable, driven by
 surface and BL top buoyancy
 fluxes, shear
- Variable boundary layer depth
- ⇒ "intermittent" downwards mixing of momnetum in growing BL
- ⇒ dilution (concentration) of
 turbulent fluxes in thick (thin)
 BL



- Marine boundary layer thickness variable, driven by surface and BL top buoyancy fluxes, shear
- Variable boundary layer depth
- ⇒ "intermittent" downwards
 mixing of momnetum in
 growing BL
- ⇒ dilution (concentration) of
 turbulent fluxes in thick (thin)
 BL





Simple model of BL variability

$$\frac{d}{dt}h = -\frac{1}{\tau_h}h + w_e^* + \frac{\zeta}{\tau_h}$$



Simple model of BL variability

$$\frac{d}{dt}h = -\frac{1}{\tau_h}h + w_e^* + \frac{\zeta}{\tau_h}$$
 vert red noise

where $\zeta \sim$ red noise



Simple model of BL variability

$$\frac{d}{dt}h = -\frac{1}{\tau_h}h + w_e^* + \frac{\zeta}{\tau_h}$$
 where $\zeta \sim {\rm red \ noise}$

Entrainment rate in momentum budget given by

$$w_e = w_e^* + \frac{1}{\tau_h} \max(\zeta, 0)$$



Simple model of BL variability

$$\frac{d}{dt}h = -\frac{1}{\tau_h}h + w_e^* + \frac{\zeta}{\tau_h}$$
 where $\zeta \sim {\rm red}~{\rm noise}$

Entrainment rate in momentum budget given by

$$w_e = w_e^* + \frac{1}{\tau_h} \max(\zeta, 0)$$

"Large-scale" winds equal u in BL

$$\mathbf{U}(z,t) = \mathbf{u}(t) \ z < h$$

and relax toward specified shear in free atmosphere

$$\frac{\partial}{\partial t}\mathbf{U}(z,t) = \frac{1}{\tau_{env}}(\mathbf{U}_s + \mathbf{\Lambda} z - \mathbf{U}) \ z > h$$



The pdf of Sea Surface Winds: Effects of Variable Stratification and Boundary Layer Height – p. 15/18

Wind speed pdf: variable BL effects





The pdf of Sea Surface Winds: Effects of Variable Stratification and Boundary Layer Height – p. 16/18

Wind speed pdf: variable BL effects





Idealised mechanistic model accounts for shape of p(w), but unable to capture large positive skewness



- Idealised mechanistic model accounts for shape of p(w), but unable to capture large positive skewness
- Surface stratification \Rightarrow insignificant influence on shape of p(w)



- Idealised mechanistic model accounts for shape of p(w), but unable to capture large positive skewness
- Surface stratification \Rightarrow insignificant influence on shape of p(w)
- "Observational" evidence that positive skew(w) associated with highly variable BL height



- Idealised mechanistic model accounts for shape of p(w), but unable to capture large positive skewness
- Surface stratification \Rightarrow insignificant influence on shape of p(w)
- "Observational" evidence that positive skew(w) associated with highly variable BL height
- Accounting in model for variable $h \Rightarrow \text{larger skew}(w)$ values, better agreement with observations



- Idealised mechanistic model accounts for shape of p(w), but unable to capture large positive skewness
- Surface stratification \Rightarrow insignificant influence on shape of p(w)
- "Observational" evidence that positive skew(w) associated with highly variable BL height
- Accounting in model for variable $h \Rightarrow \text{larger skew}(w)$ values, better agreement with observations
- Main cause: dilution (concentration) of turbulent momentum fluxes in thick (thin) BL



- Idealised mechanistic model accounts for shape of p(w), but unable to capture large positive skewness
- Surface stratification \Rightarrow insignificant influence on shape of p(w)
- "Observational" evidence that positive skew(w) associated with highly variable BL height
- Accounting in model for variable $h \Rightarrow \text{larger skew}(w)$ values, better agreement with observations
- Main cause: dilution (concentration) of turbulent momentum fluxes in thick (thin) BL
- Agreement still imperfect



- Idealised mechanistic model accounts for shape of p(w), but unable to capture large positive skewness
- Surface stratification \Rightarrow insignificant influence on shape of p(w)
- "Observational" evidence that positive skew(w) associated with highly variable BL height
- Accounting in model for variable $h \Rightarrow \text{larger skew}(w)$ values, better agreement with observations
- Main cause: dilution (concentration) of turbulent momentum fluxes in thick (thin) BL
- Agreement still imperfect
 - very coarse representation of BL processes



- Idealised mechanistic model accounts for shape of p(w), but unable to capture large positive skewness
- Surface stratification \Rightarrow insignificant influence on shape of p(w)
- "Observational" evidence that positive skew(w) associated with highly variable BL height
- Accounting in model for variable $h \Rightarrow \text{larger skew}(w)$ values, better agreement with observations
- Main cause: dilution (concentration) of turbulent momentum fluxes in thick (thin) BL
- Agreement still imperfect
 - very coarse representation of BL processes
 - model neglects feedback of winds on BL evolution



- Idealised mechanistic model accounts for shape of p(w), but unable to capture large positive skewness
- Surface stratification \Rightarrow insignificant influence on shape of p(w)
- "Observational" evidence that positive skew(w) associated with highly variable BL height
- Accounting in model for variable $h \Rightarrow \text{larger skew}(w)$ values, better agreement with observations
- Main cause: dilution (concentration) of turbulent momentum fluxes in thick (thin) BL
- Agreement still imperfect
 - very coarse representation of BL processes
 - model neglects feedback of winds on BL evolution

Need for more complex BL model; goal of future research